

Review

## Biome-Scale Forest Properties in Amazonia Based on Field and Satellite Observations

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**Abstract:** Amazonian forests are extremely heterogeneous at different spatial scales. This review intends to present the large-scale patterns of the ecosystem properties of Amazonia, and focuses on two parts of the main components of the net primary production: the long-lived carbon pools (wood) and short-lived pools (leaves). First, the focus is on forest biophysical properties, and secondly, on the macro-scale leaf phenological patterns of these forests, looking at field measurements and bringing into discussion the recent findings derived from remote sensing dataset. Finally, I discuss the results of the three major droughts that hit Amazonia in the last 15 years. The panorama that emerges from this review suggests that slow growing forests in central and eastern Amazonia, where soils are poorer, have significantly higher above ground biomass and higher wood density, trees are higher and present lower proportions of large-leaved species than stands in northwest and southwest Amazonia. However, the opposite pattern is observed in relation to forest productivity and dynamism, which is higher in western Amazonia than in central and eastern forests. The spatial patterns on leaf phenology across Amazonia are less marked. Field data from different forest formations showed that new leaf production can be unrelated to climate seasonality, timed with radiation, timed with rainfall and/or river levels. Oppositely, satellite images exhibited a large-scale synchronized peak in new leaf production during the dry season. Satellite data and field measurements bring contrasting results for the 2005 drought. Discussions on data processing and filtering, aerosols effects and a combined analysis with field and satellite images are presented. It is suggested that to improve the understanding of the large-scale patterns on Amazonian forests, integrative analyses that combine new technologies in remote sensing and long-term field ecological data are imperative.

**Keywords:** Amazonia; above ground biomass; fire; drought; field data; forest productivity; monitoring; net primary production; rain forest; satellite data; wood density; remote sensing; MODIS

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## 1. Introduction

Amazonia plays a major and yet poorly understood function in the global carbon cycle. The Amazonian tropical forest holds about  $120 \pm 30$  Pg of carbon in biomass [1] with an approximate spatial distribution of above ground live biomass ranging from  $300 \text{ Mg}\cdot\text{ha}^{-1}$  in Central Amazonia and in regions to the east and north, to  $100\text{--}200 \text{ Mg}\cdot\text{ha}^{-1}$  in the transitional and seasonal forests at the southern and north-western edges of the basin [2]. In addition, it contributes approximately 10% (4–6 Pg·C) of the world's terrestrial annual net primary productivity (NPP, the net amount of carbon that is fixed from the atmosphere into new organic matter per unit time) [3,4].

The carbon cycle has been receiving increased attention, not only from environmental scientists but also from the public and media. It has become part of the political agenda because 60% of observed global warming is attributable to the increase in carbon dioxide concentration in the atmosphere [5]. The idea that terrestrial vegetation may significantly influence the climate at regional to global scales has advanced mainly in the past three decades. Progress in technology and an increase in scientific multinational and trans-disciplinary integrative collaborations have enabled vegetation feedbacks on atmospheric and climatic processes to be studied over a wide range of scales [6,7]. For example, evaporation, condensation, cloud formation and cloud condensation nuclei particles over Amazonia have been found to play a significant role on the global circulation, affecting not only South American rainfall but also the climate of the North Atlantic and Western Europe [8–10]. On the other hand, regional patterns of rainfall, radiation and temperature act on the physiological functioning, species composition and spatial distribution of plants in the Amazon region. Studies have shown that length of the dry season in Amazonia has a significant correlation with the above ground live biomass distribution [2] and also has a major effect on geographical variation in tree community composition [11]. Water availability may also exert a yet unrecognized selective pressure on leaf shape of rainforest trees [12], and if the environmental contribution is removed (e.g., site growing conditions, soil fertility and temperature), the total annual rainfall was found to positively influence the *per area* leaf mass and some leaf nutrient concentrations [13]. Photosynthesis and productivity are stimulated by diffuse radiation, which can penetrate deeper than direct radiation into closed canopies [14]. Recent results have shown that increased diffuse radiation, caused by the increase of aerosols in the Amazon atmosphere during the dry season, increases forest  $\text{CO}_2$  uptake [15].

Despite these studies, Amazon ecosystem processes cannot be understood in terms of climate alone. It has become clear that Amazon forest structure and dynamics are strongly related to physical and chemical edaphic conditions [16]. However, elevated logistic and research costs and the high heterogeneity of these forests make a detailed understanding of these ecological observations very challenging.

Therefore, in this paper I explore the recent findings on the large-scale patterns of the forest ecosystems properties in Amazonia. First, I focus on forest biophysical properties that directly

influence the carbon dynamics across the Amazon (wood density, above ground biomass, tree height, crown size and wood productivity). Then, I explore the macro-scale phenological patterns of these forests, and bring into discussion the recent findings derived from remote sensing dataset. Finally, I present the most recent findings on the effects of droughts in Amazonia. A review of climate change effects on tropical forests have been explored elsewhere [17].

## 2. Forest Biophysical Properties

### 2.1. Wood Density

Wood density is an important factor in converting forest volume data to biomass. It is defined by the oven-dry weight divided by wet volume [18]. Wood specific gravity is highly correlated with the density of carbon per unit volume and is thus important for estimating ecosystem carbon storage and fluxes [18,19]. However, one of the large sources of uncertainty in all estimates of carbon stocks in tropical forests is the lack of information on tree height and wood density due to the high number of species [20].

The distribution of wood specific gravity within a tropical tree community is theoretically expected to vary among sites, associated with one or a combination of soil fertility, rates of forest disturbance, early and secondary successional vegetation, tree growth and mortality [21–26]. High disturbance rates and high turnover rates are expected to occur in faster-growing species, which as a result have relatively low density woods, while low soil fertility slows tree growth, and is expected to lead to relatively high density woods [27]. However, [24] argues that there are no clear explanations for differences in wood specific gravity among sites, and it seems likely that different factors may act at different scales.

In Amazonia, it has been suggested that regional patterns of species composition and abundance drive the observed east-west gradient in wood specific gravity [28]. Overall, mean stand-level wood specific gravity is 16% higher in forests in central and eastern, compared with north/western Amazonia [27]. It has been suggested that the higher specific gravity in central and eastern Amazon forests is related to the regular seasonal water availability or El Niño related droughts that occur in this region [23], and more recently, to a certain extent to the soil physical structure and quality [16].

### 2.2. Above Ground Biomass

A number of approaches have been used to estimate the spatial distribution of biomass in Amazonia, ranging from the development of allometric equations [25,29,30], local and large scale forest inventories [1,28,31–34] to a methodology using only remote sensing or combined with field data [2,35–37]. Houghton *et al.* [38] compared several biomass estimates for the Brazilian Amazon, and found very low agreement across estimates.

Differences in biomass estimates based on field measurements arise from differences in measurement methods in individual plots, extrapolation of the data from individual plots for the region, and the heterogeneity of the forests [38,39]. Remote sensing-based estimates are usually restricted either to a small area of coverage with high spatial resolution (e.g., Landsat, IKONOS, Japanese Earth Resources Satellite (JERS) or airborne laser data ([40,41] and others) or else to sensors

with a higher area of coverage, but with low spatial resolution [42] which implies a loss of accuracy to detect variations in forest physiognomy. Direct estimates of biomass based solely on optical or radar data derived from satellites present inconsistencies in forests with moderate and high biomass, and new technologies are being developed, such as the BIOMASS mission from the European Space Agency (ESA) [43]. Recent studies based entirely on field data [1] or on field data merged with a remote sensing dataset [2], suggest high biomass ( $300\text{--}400\text{ Mg}\cdot\text{ha}^{-1}$ ) in north-eastern Amazonia and central areas west of the Trombetas river to the west of the Rio Negro river, lower biomass ( $250\text{--}300\text{ Mg}\cdot\text{ha}^{-1}$ ) in the main channel of the Amazon River, Igapó and Várzea floodplains and between  $200\text{--}300\text{ Mg}\cdot\text{ha}^{-1}$  in western and southern Amazonia (Figure 1(a,b)). At a global scale, AGB has been recently estimated by using a combination of remote sensing and field data: the global forest height data measured by the Geoscience Laser Altimeter System (GLAS), onboard the Ice, Cloud, and land Elevation Satellite (ICESat), the moderate resolution imaging spectroradiometer (MODIS), the shuttle radar topography mission (SRTM), and quick scatterometer (QSCAT) [44]. The comparison of this global AGB with the map produced by [2] revealed in this new dataset lower biomass in the Rio Negro Basin, central region, and higher biomass in western Amazonia. Western Amazonia is a region that presents lower biomass than central regions [1,23,27,45] and it could be speculated that the bamboo forests, present in this region, could have effects on Lidar measurements.

**Figure 1.** Biomass and canopy height estimates for Amazonia. (a) Interpolations of biomass calculated by overlaying basal area estimate with maps of the structure and wood density function (Adapted from [1]); (b) 1 km spatial resolution map derived from combined remote sensing and field data through a decision tree method and regression analysis with 11 biomass classes and overall accuracy of 88% (Adapted from [2]); (c) 1 km spatial resolution canopy height data from [53], available online: <http://lidarradar.jpl.nasa.gov>.

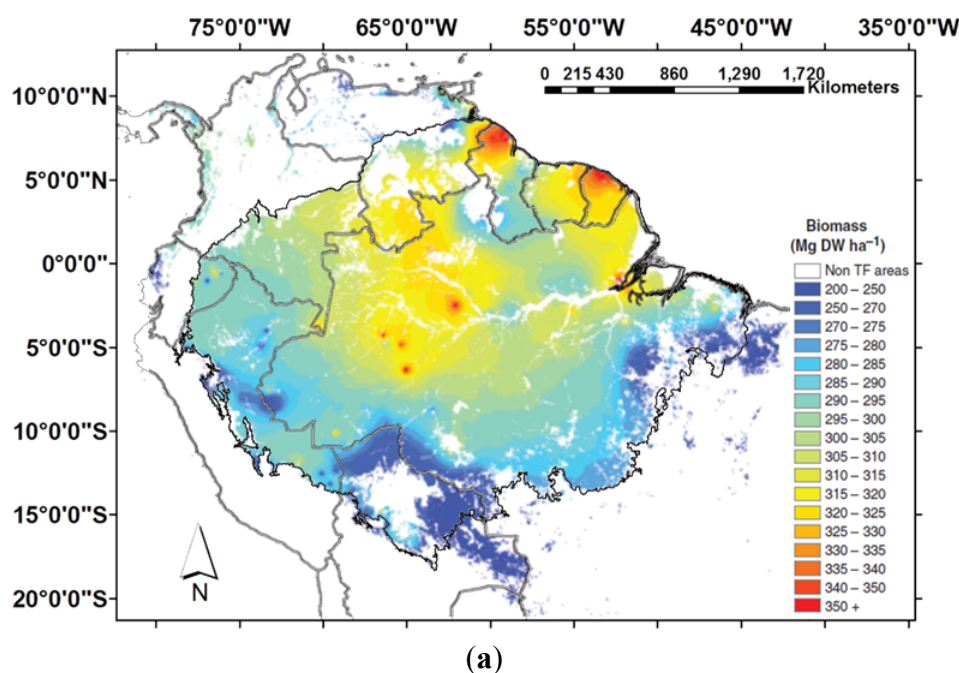
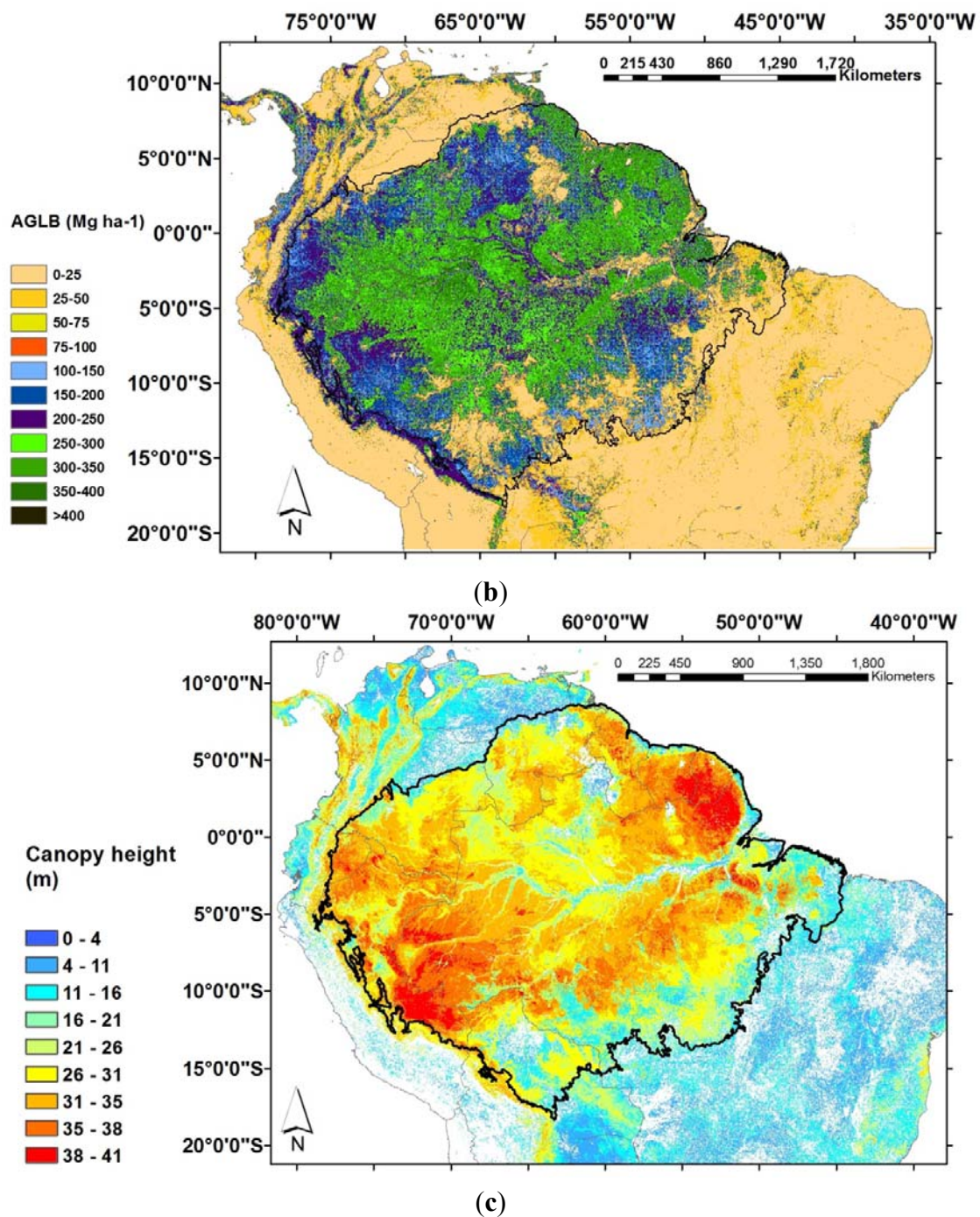


Figure 1. Cont.



The use of RADAR for biomass estimation in dense canopy forests is still limited. At a local scale, Synthetic Aperture Radar (SAR) has been used to map and estimate biomass on different land covers. By analyzing SIR-C data in C- and L-bands in a region in central Brazilian Amazonia, [46] noted a visual discrimination of the main regenerating classes, indicating a potentiality for biomass estimation, although the analysis of the six bands separately produced no significant relationships between SAR backscatter and forest biomass. The analysis performed by [47] using JERS-1 L-band data for central Amazonia suggested that only three broad classes of regenerating forest biomass density were distinguished and saturation were observed at lower biomass when compared with saturation limits in



other forest types such as a coniferous plantation. Improvements on biomass mapping were achieved by the use of airborne P-band SAR [48], although primary and secondary forests in some cases exhibited similar P-band backscatter values. For the same region as [48], Gonçalves *et al.* [49] found average prediction errors of less than 14% for quantifying stem volume by using polarimetric synthetic aperture radar (PolSAR). A review on mapping and monitoring biomass estimation by comparing different methods and satellite data is provided by [50].

### 2.3. Canopy Height

There are few studies focusing on the spatial variability of tree height in Amazonia. The variation in tree height (and thus wood volume) with environmental factors are also poorly described. Nonetheless, tree height is expected to show a similar pattern to basal area and decrease with increasing dry season length as hydraulic constraints on tree height become more severe. Moreover, height increase in large diameter trees is expected to be limited due to mechanical and physiological constraints [51].

The development of a new approach to map forest height globally using light detection and ranging (Lidar) data from the Geoscience Laser Altimeter System (GLAS) is expected to greatly contribute to the understanding of forest structure, and refinements in the spatial scale would provide critical information on the pools' above-ground biomass. The GLAS products were used by [52,53] to generate a global canopy height map. Both maps detected higher errors in closed broadleaves forests, such as the Amazon. In the Amazon basin Simard *et al.* [53] found differences in canopy height as a function of distance from rivers and edge effects were observed along roads and in the arc of deforestation (Figure 1(c)).

At local scale, the use of the Laser Vegetation Imaging Sensor (LVIS) detected canopy top height changes between the data from 1998 and 2005 in secondary and old growth forests coincident to the expected changes of these land cover types [54].

### 2.4. Crown Size and Leaf Size Patterns

A macrostructure of crown size has been detected by [55] in Amazonia *terra firme* forests. Despite the high local variability, apparent crown size tends to be larger than 16 m in a broad east-west band, from 5–10° of the equator, and larger than 15 m in the contiguous region in north-eastern Amazonia. These areas of large tree crown size are broadly located in the regions with short dry season (1–3 months). On the other hand, small crown sizes (<14 m) dominated north-western and the very southern limits, close to the Cerrado biome. These areas correspond to regions with no dry season or long dry season (>5 months). Interestingly, north-west Amazonia presents significantly greatest proportions of large-leaved species, and is followed by southwest, central and eastern and finally, north Amazonia [56]. In addition, species with large leaves have lower wood density (approximately 10% less) when compared with species with small leaves [56].

### 2.5. Forest Productivity

The biochemical construction of new organic material over a specified time interval—known as net primary production (NPP)—of tropical forests and its partitioning between long-lived carbon pools

(wood) and shorter-lived pools (leaves, fine roots) is of considerable importance in the global carbon cycle [4]. It has been suggested that the NPP as well as the nitrogen and phosphorus content, all scale allometrically with phytomass across diverse plant communities, from tropical forest to arctic tundra [57].

The quantification of NPP is still very uncertain but considerable work has been done on the assessment of the above-ground components of NPP (leaf, flower, fruit and wood production) for different ecosystems. On the other hand, below-ground NPP, the partitioning between above- and below-ground components and the main environmental drivers of these patterns are still poorly understood. Data on both below- and above-ground NPP are especially limited in Amazonia due to elevated costs and challenging logistics involved in setting up and monitoring measurement sites. There are different approaches to evaluate and quantify the below- and above- ground primary production. Field measurements for assessing the above-ground component should include litterfall, branch production, coarse woody biomass production, aboveground biomass losses to consumers, and the emission of volatile organic carbon compounds and aboveground losses of leached organic compounds [58–62]. For the below-ground component, data on coarse and fine root productivity, dead coarse and fine roots, losses to consumers, and the carbon loss through exudates and mycorrhizae are necessary [61,63]. The second option for evaluating the above-ground component (mainly canopy process) relies on the use of remote sensing data and modeling [64–68], although field measurements are important for validation purposes.

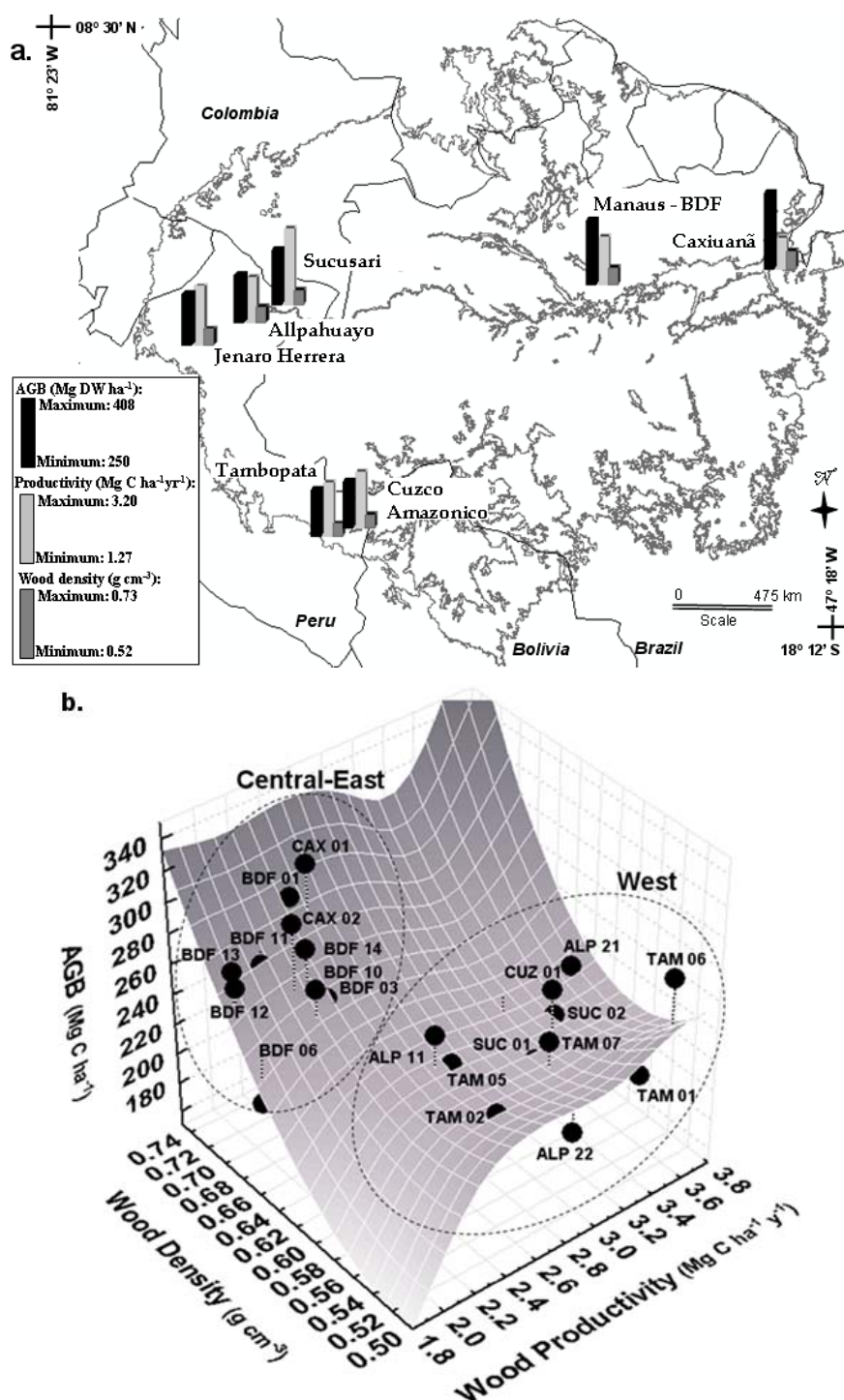
Looking at the above-ground coarse wood carbon productivity in stems and branches, Malhi *et al.* [60] demonstrated that wood production varies by up to a factor of three across Amazonian forests. Lower wood production, according to that study, is found on poor oxisols in lowland eastern Amazonia and higher production occurs in more fertile soils in western Amazonia. In relation to the canopy NPP (leaves, twigs, flowers and fruits), Chave *et al.* [62] suggested that soil is not a major determinant of litterfall patterns, however infertile white sandy soils have significantly lower litterfall than other soil types, and more fertile soils exhibit higher canopy production [61]. Below-ground allocation declines with increasing clay content and more fertile soils tend to have higher fine root production [61]. Overall, soil phosphorus is likely to be more important than nitrogen in the determination of the total NPP across the Amazon [16]. It has also been suggested that forests in infertile soils, such as eastern Amazonia, have low carbon use efficiency, in comparison with recently disturbed forests or forests in fertile soils [69].

Remote sensing and modeling studies suggest or assume that temperature followed by solar radiation primarily determine variation in NPP in Amazonia, which is especially sensitive to large ENSO events [70,71] although soil depths significantly affect the gross primary production (GPP, the sum of net primary production (NPP) and autotrophic respiration) [68,72]. While temperature appears to exert a strong influence on NPP in some models [64], other models suggest that radiation increases carbon fluxes in Amazonian [65,73]. This is also suggested by remote sensing observations, which showed increases in leaf area index (LAI) and Enhanced Vegetation Indices (EVI) during the dry-season [66,74]. It is important to mention that not only remote sensing data but also modeling studies have limitations when applied to evergreen canopies such as the tropical forests, due to saturation in the data in the first case and lack of information in the latter.

A synthesis of the large-scale patterns of wood density, above ground biomass and wood productivity, compiled from [1,27,60] is presented in Figure 2. By combining these biophysical

measurements, two distinct groups of plots emerge: “east-central”, with higher above ground biomass, wood specific density and lower productivity than the “west” group (Figure 2(b)). A detailed evaluation of the plots and field sites is presented in [45].

**Figure 2.** Spatial distribution of above ground biomass, wood productivity and wood density across Amazonia: (a) Average values per site; (b) Plot values. Data from [45].



Although there is a great effort from the scientific community on developing new techniques for mapping forests biomass, the understanding of how these forests will change with climatic variations, not only in terms of stocks but also in terms of metabolism or productivity represents perhaps a greater



challenge. The differences among the biomass maps developed by [1,2,44] are possibly minor if compared with the uncertainties in those estimates. The availability of new products, such as the global carbon map and the global tree height map, [44] and [53], respectively, will certainly produce a strong “footprint” on the current state of the world’s forests. The development of such dataset for the next years will be important for the monitoring point of view, but remains uncertain if changes in biomass will be detectable with the available methods and at large scales.

In summary, by dividing the western Amazonia in north and south regions, we can conclude that north-western Amazonian forests presents lower wood specific gravity, small tree crown, greater proportions of large-leaved tree species, tree height varying from 26 to 38 meters, higher productivity and lower above ground biomass than the eastern regions. South-western forests differs from north western forests by presenting relatively lower biomass, lower proportion of large-leaved species and higher tree height. Eastern forests present higher wood density and higher above ground biomass, lower productivity, large tree crown, lowest proportions of large-leaved species and canopy height from 26 to 41 meters.

### 3. Vegetation Phenology

The term *phenology* is derived from Greek and means ‘to show or to appear’. In the ecological context, phenology is defined as the study of the seasonal timing of life cycle events [75], which can be critical for survival and reproduction of plants. Moreover, phenological characteristics of species reflect the influence of evolution and the environment on plant traits. These traits in turn have substantial implications for plant functioning at the level of the leaf, plant and ecosystem.

A complex mosaic of vegetation phenology is expected for the highly heterogeneous Amazonian forests as a result of differences in forest structure and composition associated with geomorphological and pedological conditions as well as variation in rainfall and radiation regimes.

In the next two sub-sections, I outline some evidence and discuss the phenological timing of canopy processes in Amazonia through the conceptual lens of field ecology, and the more recent insights from the remote sensing community. More specifically, the focus is on leaf phenology, as its periodicity, synchronicity and variability are still poorly understood and quantified.

#### 3.1. Phenological Patterns Observed on the Ground

Evergreen, deciduous and semi-deciduous tropical tree species have leaves with different life spans. Some (not mutually exclusive) hypotheses that could explain how this cycle relates to environmental factors could be [76]: (a) competition for nutrients controls the senescence of individual leaves; (b) deciduous leaves evolved as a drought-avoidance mechanism; (c) the trade-off between leaf costs and benefits controls the life span; and (d) long-lived leaves have a high nutrient use efficiency (defined as total plant dry mass divided by foliar nitrogen or phosphorus mass). Aide [77] suggested that at a broader scale, some biotic factors could act on the leaf phenological cycle, and demonstrated that herbivores could influence leaf phenology of understory plants in a tropical forest in Panama, as 34% of the annual leaf production was observed during the dry season when insect abundance and damage to expanding leaves were lower.

The vegetation in the floodplains along the rivers in Amazonia is exposed to flood pulse hydrological regimes and experiences periodical inundation of large areas of forest. In these areas, there are two main vegetation formations: *Várzeas* and *Igapós*. *Várzeas* can be defined as forests flooded by white-water rivers (sediment rich rivers originating in the Andes), while *Igapós* are nutrient-poor areas flooded by black water Rivers (sediment poor, organic matter rich rivers originated in lowland forests) [78]. Despite the variability of phenological traits in these physiognomies, studies in central Amazonia suggest that the flush of new leaves in *Várzea* occurs in August in most species, while in the *Igapó* vegetation type, new leaves are flushed in July and August, and also in October, December and February [79–84].

Generalizations for temporal pattern of leaf production in *Terra Firme* forests are more difficult to perceive due to the lack of one strong event to mark or homogenize a signal to correlate with the vegetation response, such as the flood pulses in the forests along the rivers. Regular periods of low rainfall could act as such an event but unfortunately there is not a standardized way to define dry season in the literature, making more difficult an integrative analyses. For instance, the dry season definition can vary from the period of lower rainfall without a defined threshold ([73] and many others) to a very clear defined threshold, such rainfall less than  $50 \text{ mm} \cdot \text{month}^{-1}$  [85] or less than  $100 \text{ mm}$  rain per month [86]. It has been suggested that dry season can be characterized by the periods during which evapotranspiration exceeds precipitation and the soil water available to plants declines [87] and  $100 \text{ mm} \cdot \text{month}^{-1}$  rainfall could be used as a mean modeled threshold for Amazon evergreen forests [88].

In addition to temporal patterns of rainfall, seasonal variation in irradiance also occurs because of changes in cloud cover between dry and wet seasons and solar angle throughout the year and in different latitudes, affecting tropical forest canopies [13,89]. In a review of the literature, [90] suggested that tropical tree species could take advantage of maximum irradiance to produce new organs, conferring on them two main benefits (although co-varying environmental factors cannot be excluded). First, studies have shown that there is an increase of phenological activity and net primary production during the months of solar radiation peak when water is available. Secondly, leaf flush during the dry season would represent an advantage as insect activity and fungal pathogens are reduced during this period. More recent studies also support these hypotheses [14,91,92], although bimodal distributions of flushing of new leaves have been reported [91,93,94] as well as regions with negligible relationships between leaf production and leaf mortality with climate seasonality [95] and species with new leaves production during the wet season [96]. A recent review of litterfall seasonality in Amazonian forests demonstrated that it does not depend on annual rainfall across different sites in old growth forests nor on soil type [62]. However, their results showed significant positive relationships between litterfall seasonality and rainfall seasonality and higher litterfall production in forests growing on fertile soils than in forests growing in infertile soils in Amazonia.

In conclusion, the variability of leaf phenology in the Amazon is still poorly understood due to the high heterogeneity of forests types, the relative paucity of observational descriptions in the literature for such a large region and difficulties in inter-comparisons between events in the dry or wet seasons due to inconsistent definitions of seasons. Therefore, in the next section I explore how remote sensing data have been used to explore the temporal variation of phenological process covering large areas in Amazonia.

### 3.2. Phenological Patterns Derived from Remote Sensing

#### 3.2.1. Background

Remote sensing technology offers a possible solution to the problem of collecting leaf data from the largest area of continuous forest on Earth. However, a potential obstacle to the acquisition of a multitemporal dataset for detecting vegetation phenology in the Amazon by satellite-derived vegetation indices is the influence of aerosols from biomass-burning [97,98] and clouds and cloud's shades [99]. These problems are slowly becoming resolved through the use of a new generation of moderate-resolution sensors with more frequent coverage (e.g., daily) of Amazonia. These sensors represent a new opportunity for frequent and efficient assessment of natural changes in forest canopies. For instance, the Moderate Resolution Imaging Spectroradiometer (MODIS) instruments aboard the Terra (morning) and Aqua (afternoon) platforms provide consistent daily coverage of the entire globe at 250 m to 1 km resolution with 36 bands of spectral information since 2000 and 2001, respectively [100]. The MODIS instruments acquire images up to 4 times a day, with a swath width of approximately 2300 km. The data are geometrically and radiometrically rectified, and land-cover products are automatically generated for the monitoring of the earth surface, such as the daily and 8-day composite Surface Reflectance product (MOD09), 16-day composite Surface Reflectance and Vegetation Index product (MOD13), 8-day composite Leaf Area Index (MOD15), Vegetation Cover Conversion product (MOD44A), and the Vegetation Continuous Fields product (MOD44B).

Vegetation index images are the most widely used satellite data to monitor natural vegetation dynamics as they measure canopy greenness, a composite property of canopy structure, leaf area and canopy chlorophyll content [101]. Over the last several decades, vegetation monitoring at regional and global scales has been done with the Advanced Very High Resolution Radiometer (AVHRR) sensor data converted to Normalized Difference Vegetation Index (NDVI) images [102–108]. NDVI is defined as:

$$NDVI = \frac{(\rho_{NIR} - \rho_{Red})}{(\rho_{NIR} + \rho_{Red})} \quad (1)$$

where  $\rho_{NIR}$  and  $\rho_{Red}$  are the reflectance of NIR and Red spectral bands.

In addition to the NDVI, the MODIS vegetation indices product also has the Enhanced Vegetation Index (EVI). NDVI has been intensively studied, and limitations including saturation in closed canopy and sensitivity to atmospheric aerosols and soil background have been reported [109,110]. In this sense, EVI was developed to minimize these effects, including the blue band for atmospheric correction, computed as (Equation (2)):

$$EVI = 2.5 \times \frac{\rho_{NIR} - \rho_{Red}}{\rho_{NIR} + (6 \times \rho_{Red} - 7.5 \times \rho_{Blue}) + 1} \quad (2)$$

where  $\rho_{NIR}$ ,  $\rho_{Red}$  and  $\rho_{Blue}$  are the reflectance in the NIR, Red and Blue channels respectively; 2.5 is a gain factor, 6 and 7.5 are coefficients designed to correct for aerosol scattering and absorption and 1 is a canopy background adjustment [111,112]. While NDVI is more sensitive to the absorbing chlorophyll bands (red) and to atmospheric properties such as clouds, water vapor solar illumination and satellite viewing geometry [113,114], EVI is more sensitive to canopy structural variations, including leaf area index [111].

### 3.2.2. Monitoring Forest Phenology

The first global dataset with high temporal resolution was acquired by the Advanced Very High Resolution Radiometer (AVHRR) sensor, onboard the TIROS-N satellite (launched in 1978) and National Oceanic and Atmospheric Administration (NOAA) series 6 through 14 (the latter launched in 1994) meteorological satellites. A number of vegetation indices were developed based on AVHRR images, including: the Normalized Difference Vegetation Index (NDVI) [115–118], the Vegetation Conditioning Index (VCI) [118,119], the Temperature Conditioning Index (TCI) [119,120], the Vegetation Health Index (VHI) [121], the NDVI-Land Surface Temperature (NDVI-LST) [122], the Drought Severity Index (DSI) and the Palmer Drought Severity Index (PDSI) [123–125].

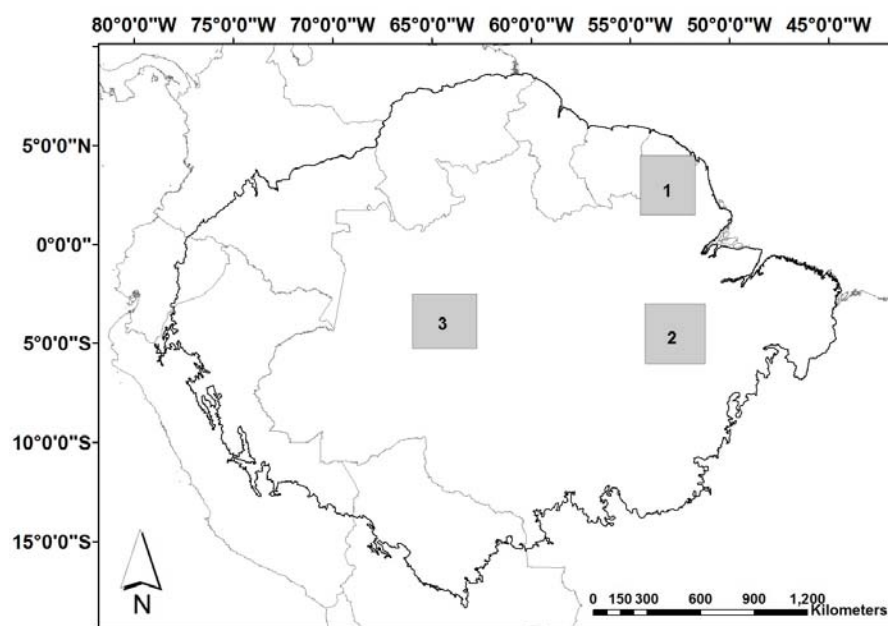
Then, a more specific water sensitive index was developed, the normalized difference water index (NDWI) [126]. This index showed adequate potential for canopy-level water content estimation ([127–129] and others).

One of the first studies for the monitoring of Amazonian forests was carried out by [130]. Using daily NDVI images derived from NOAA AVHRR from 1982 to 1988, [130] found little or no cyclical pattern over the Amazon forest. They also reported a decline in the response of the vegetation index over the dense forested area over time, demonstrating the effects of the sensor degradation and the lack of sensor calibration. Liu et Kogan. [131] studied drought patterns in South America from 1985 to 1992 using AVHRR data. They found that NDVI and VCI agreed well with rainfall anomalies for the entire country. However, the relationship between NDVI and monthly rainfall in the tropical forest was weak and the drought dynamic patterns for both indices in this region were not strongly correlated with the meteorological precipitation anomaly data. Xiao *et al.* [132] carried out a correlation analysis between NOAA AVHRR NDVI and field meteorological stations in the Amazon evergreen forest, finding no relationships. However, during strong climatic events, such as El Niño and La Niña, changes in the amplitude of AVHRR NDVI over Amazonia suggest that the forest phenology is responsive to rainfall variation over seasonal and inter-annual time scales [106]. It has also been suggested that temporal variations in southern and eastern forests of Amazonia in the NOAA AVHRR NDVI data could be primarily associated with atmospheric variations such as clouds and aerosols [98].

The MODIS sensor, which has improved radiometric and atmospheric corrections in relation to the NDVI generated by the AVHRR sensor, has greatly contributed to the advances in the knowledge of the Amazon forest phenology. Studies evaluating the EVI for the Amazon evergreen forest captured for the first time a remarkable large-scale seasonal pattern. Huete *et al.* [74] observed an increase of 25% in EVI (enhanced leaves activity) in the sunnier dry season across Amazon forests, suggesting that sunlight may exert more influence than rainfall in the phenology of this forest. Results from the study performed by [133] using MODIS EVI also reinforce this result. Quantitative analysis of the MODIS LAI product for the Amazon forests demonstrated that changes in LAI are positively correlated with changes in solar radiation and negatively correlated with changes in precipitation, although the correlations between leaf area and radiation changes are larger and more numerous [66]. However, Poulter *et al.* [68] showed that the MODIS LAI seasonal differences are particularly sensitive to additional filtering for indirect aerosol and cloud effects, while the seasonal differences in EVI were less sensitive to the filtering methods evaluated.

The large scale patterns observed on the new leaf production [66,74,133] encompasses many different forest types and these results do not agree with the observation of field studies, where peak in leaf flush has been observed during the dry, wet or dry and wet seasons in different forests in Amazonia. Although Brando *et al.* [134] observed a positive correlation between EVI with leaf flushing measured in the field in central-east Amazonia, EVI was relatively insensitive to changes in LAI. The increase in EVI during the dry season has also been attributed to solar illumination effects rather than changes in LAI; and changes in canopy foliage detected from MODIS LAI data were not consistent with LAI estimates from hemispherical photographs for a site in southern Amazonia [135]. Another criticism to the results presented by [66,74,133] is the fact that areas that exhibited higher indices during the dry season in their analysis are located in regions with dissimilar onset and end of the dry/wet season [136]. For example, areas that presented higher indices in their study are located in the north hemisphere in the Amazon region—north-east Amazonian forests—and have the dry season from September to November (sample 1, Figure 3); in center-east Brazilian Amazon—the Tapajós region, low rainfall occurs from July to October (sample 2, Figure 3), and central-west Brazilian Amazon presents low rainfall in August (sample 3, Figure 3). These contrasting results show that the large-scale extrapolation of a data collected from one site and with a reasonable agreement with the satellite image may not be valid for different regions. In addition, saturation and noise in the remote sensing dataset may cover the detection of changes in such dense canopy forests.

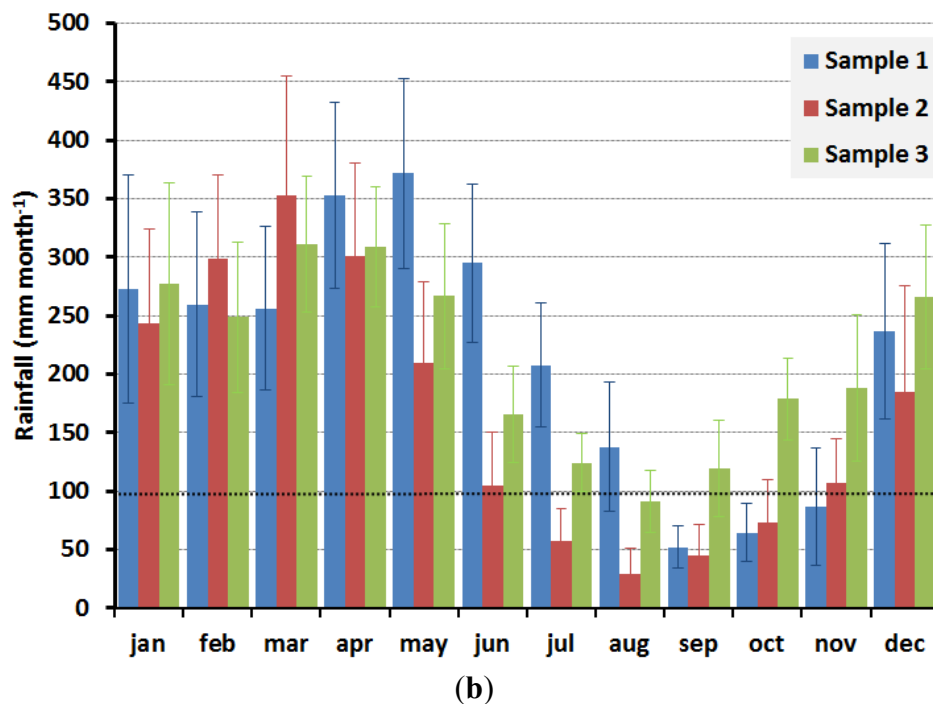
**Figure 3.** Rainfall derived from TRMM (Tropical Rainfall Measuring Mission), product 3B43-v6, at 0.25° spatial resolution. Dry season is defined here when rainfall is lower than the estimated tropical forests canopy transpiration ( $\sim 100 \text{ mm} \cdot \text{month}^{-1}$ ). (a) Location of the samples sites; (b) Averaged rainfall values for each sample, from 1998 to 2010, and the standard deviation as error bars.



(a)



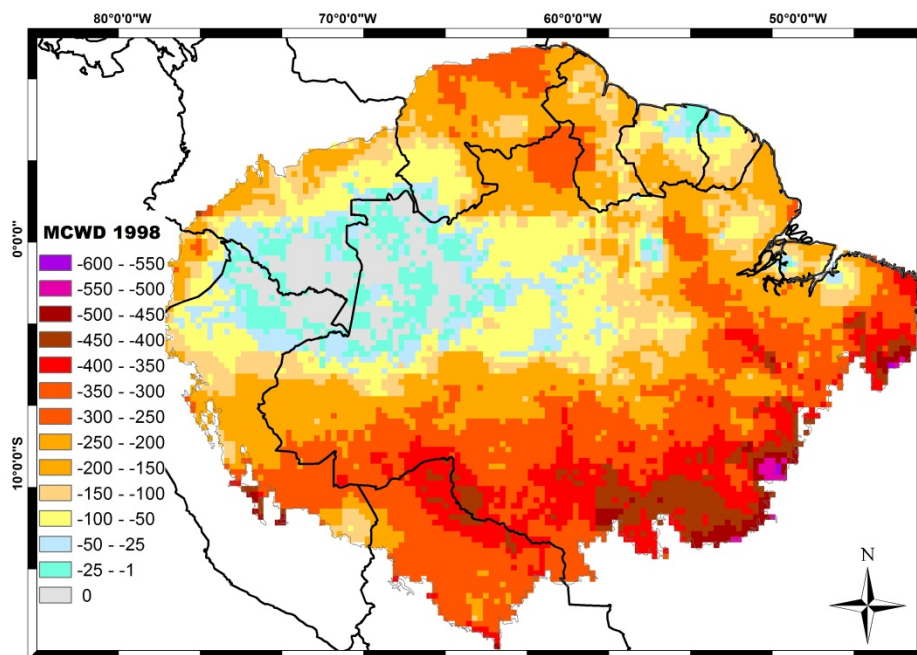
Figure 3. Cont.



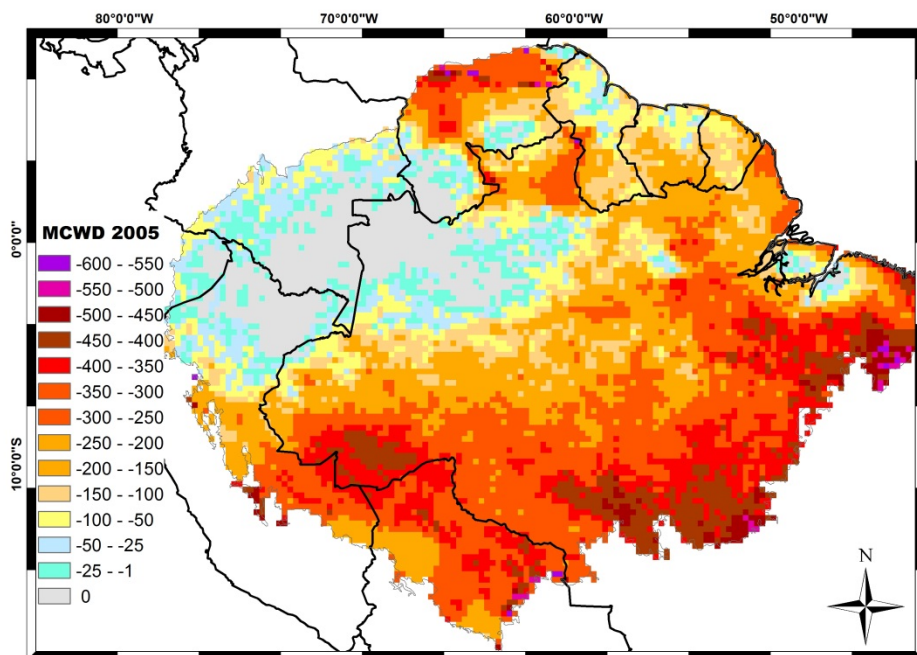
#### 4. Drought Effects

The El Niño, associated with the low phase of the Southern Oscillation (SO), coincides with reduced Amazon rainfall, particularly in the northern and central regions. Opposite anomalies often occur during the high phase of the SO, related to La Niña. The 1997/98 drought during an El Niño year was considered the strongest in that century [137]. Differently from the El Niño-related droughts, during 2005, large sections of southwestern Amazonia were severely affected, and this drought was considered one of the most intense droughts of the last hundred years. The 2005 drought was driven by the warming of the tropical North Atlantic, combined with the reduced intensity in northeast trade wind moisture transport into southern Amazonia and the weakened upward motion over this section of Amazonia [138]. Five years later, another strong drought hit the Amazon. The 2010 drought started during an El Niño event in early austral summer of 2010 and then became more intense during La Niña in the austral winter dry season and the following spring. These three recent droughts affect different regions of the Amazon (Figure 4). In 1998, southwest regions of the Amazon, such as Bolivia and Rondonia state in Brazil, and north-eastern regions such as Roraima state were deeply affected. In 2005, the MAP region (Madre de Dios in Peru, Acre in Brazil and Pando in Bolivia) were the most severe affect region. In 2010, both central-north and central-south Amazonia were the most affected areas, covering Roraima and Mato Grosso states in Brazil and Guiana.

**Figure 4.** Drought severity can be measured by the maximum cumulative water deficit (MCWD), which corresponds to the maximum value of the accumulated water deficit (WD) that reached each area within a period of time. The WD is estimated based on the approximation that moist tropical forests transpires  $100 \text{ mm} \cdot \text{month}^{-1}$ , and therefore, when rainfall is lower than  $100 \text{ mm} \cdot \text{month}^{-1}$ , the forest is considered under water stress. Areas that experienced more severe droughts present lower MCWD values. Spatial distribution of the MCWD between January and December (a) 1998; (b) 2005; and (c) 2010. Data derived from TRMM (Tropical Rainfall Measuring Mission), 3B43-v6, at  $0.25^\circ$  spatial resolution.

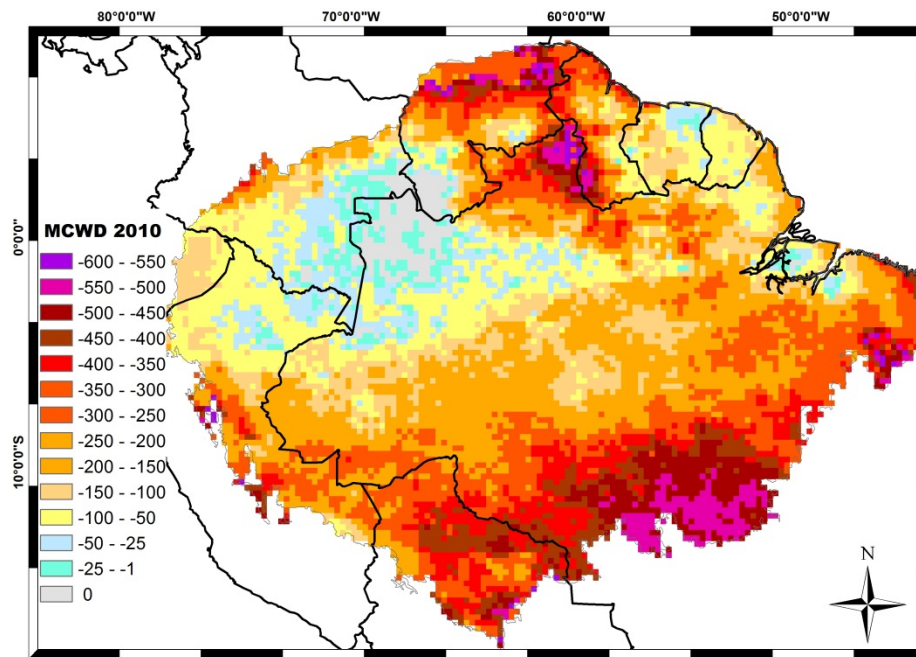


(a)



(b)

Figure 4. Cont.



(c)

Long term measurements in central Amazonia showed that the 97/98 ENSO drought increased the tree mortality (1.91%) in comparison with 5- to 13-year measurements average (1.12%), and that trees that died due to the drought did not differ significantly in size or species composition from those that died previously [139].

After the 2005 drought extreme, the first study published was carried out by [140]. They evaluated the anomalies in the EVI generated from the MODIS c4 collection during the driest quarter of 2005 and found a significant increase in the EVI during the peak, rather than a decline. This was interpreted as a possible increase in productivity during the drought period. However, two years later a study based on field measurements evaluating the impact of the 2005 drought demonstrated that plots most affected by the rainfall suppression exhibited a decline in the rate of net aboveground biomass accumulation, and those losses were driven by occasionally large mortality increases and by widespread decline in growth [141]. More information on the effects of droughts in the Amazon can be found in the special issue of the *New Phytologist*, volume 187, issue 3, 2011. These contrasting results intrigued the scientific community and a re-assessment of the EVI processing was carried out by [142]. The authors reported differences with respect to the area and intensity of the 2005 increase in the EVI between results obtained from the c4 and c5 product and thus contested [140] results. Anderson *et al.* [143] also evaluated the anomalies in the EVI c5, and found only small and scattered areas with increase in this index. Moreover, Anderson *et al.* [143] showed that areas with positive anomalies in EVI presented positive anomaly in tree mortality, an indication that the increase in the EVI could be related to changes in the canopy structure. This result is also supported by a previous study where it is shown that the decline in the shade fraction is highly associated with the increase in EVI [144]. There is no conclusive result that fully explains what drove the increase in EVI during the 2005 drought. Although image processing artifacts could explain part of this result [142], no relationship was found between areas with EVI anomalies and areas with anomalies in atmospheric

optical depth (a measurement of airborne particles, such as dust, cloud droplet, ice particles, biomass smoke, *etc.*) [143].

The preliminary study on the effects of the 2010 drought estimates a higher impact on the forests than the 2005 drought [145], and an evaluation of MODIS vegetation indices observed a decline in the greenness in an area four times greater (2.4 million km<sup>2</sup>), more severe than in 2005 and that persisted even after the return of normal rainfall levels, unlike in 2005 [146]. There is no study published based on field measurements yet, but a cumulative effect of the droughts could be expected.

In addition to the effects measured in primary forests, droughts also play a major role on the increase in fire susceptibility in Amazonian forests. It has been estimated that fires in the Amazon increased by 42% from 1998 to 2006 [147], with highest incidence in the arc of deforestation [148]. For instance, the 1997/98 drought affected the savannahs in Roraima, as 53% of the area burned in 1997/1998 in contrast to 1998/1999 (a wet year: 30% burned) and 1999/2000 (a normal year: 36% burned) [139]. From the 6,500 km<sup>2</sup> of land surface burnt in Acre State during the 2005 drought, 2,800 km<sup>2</sup> corresponded to areas of standing forests [149]. A great spread of fires on primary forests and forest fragments can also be expected to have occurred in 2010, as one of the most affected regions is located in the arc of deforestation. Therefore, fires can be considered an active factor for large-scale changes in forest structure and dynamics in this biome.

## 5. Future Directions

The improvements in remote sensing data and techniques, and the effort to install and monitor long-term ecological sites are necessary to answer local and regional questions. Integrative analyses are still very restricted, and probably are the next challenge to improve our knowledge of Amazonian.

Multi-collaborative projects in Amazonia have a successful history on acquiring, documenting, publishing and sharing ecological data. Examples of these programs include the Biological Dynamics of Forest Fragments Project (BDFFP), Center for Tropical Forest Science (CTFS), Tropical Ecology Assessment and Monitoring Initiative (TEAM), the Amazon Forest Inventory Network (RAINFOR), among others. However, due to the unlikely possibility of setting a randomly distributed network of field plots covering the Amazonian landscapes, remote sensing data can play a key role. For example, one step that could be adopted by scientist is the landscape evaluation of the area where they intend to set a new plot/site. Landsat data has proven to be adequate for detecting different forest physiognomies in dense Amazonian forests [45,150,151], and it has been released free of charge by the National Institute for Space Research (INPE).

A better documentation and standardization of field data and protocols is also needed. Although field courses, workshops and technical training has been shown to be highly effective, the limited resources requires a better use of the advances in the World Wide Web, such as to improve communication and data sharing (e.g., [152]) or to provide more details on site and data collections (e.g., using the on-line Supporting Information section of scientific journals).

The temporal scale for monitoring the different components of the Amazonian forests is a key issue for remote sensing data integration. Although for the quantification of above ground biomass, forest growth and recruitment, and changes in species composition can be achieved with inter-census intervals of 2 or more years, the understanding of canopy processes, such as direct observation of leaf

phenology requires intensive measurements. Monthly images derived from remote sensing datasets showed to detect changes in the canopy of densely vegetated areas of Amazonia, and the validation of these observations are necessary. Despite increasing the costs for setting a plot with monthly measurements, this information is essential for the understanding of how tropical forests are responding to changing environmental factors.

As drought frequency has been predicted to increase in the Amazon, it is also expected that fires will increase. Techniques, such as the linear un-mixing model have proved adequate to detect burn scars in primary forests. However, the development of a methodology to detect the levels of susceptibility of primary forest to fires would be a great contribution, as it could be used as an “alert system” of forest flammability. Nonetheless, to achieve this objective, more research that combines both remote sensing and field measurements over burnt areas are necessary to fully understand the local factors and conditions that triggers the spread of fires and its intensity on primary forests.

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